



1 Article

2 Effects of Bank Vegetation and Incision on Erosion 3 Rates in an Urban Stream

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- 7 Received: date; Accepted: date; Published: date
- 8 Abstract: Changing land-use associated with urbanization has resulted in shifts in riparian
- 9 assemblages, stream hydraulics, and sediment dynamics leading to the degradation of waterways.
- 10 To combat degradation, restoration and management of riparian zones is becoming increasingly
- 11 common. However, the relationship between flora, especially the influence of invasive species, on
- 12 sediment dynamics is poorly understood. Bank erosion and turbidity were monitored in the
- 13 Tookany Creek and its tributary Mill Run in the greater Philadelphia, PA region. To evaluate the
- 14 influence of the invasive species Reynoutria japonica (Japanese knotweed) on erosion, reaches were
- 15 chosen based on their riparian vegetation and degree of incision. Bank pins and turbidity loggers
- 16 were used to estimate sediment erosion. Erosion calculations based on bank pins suggest greater
- 17 erosion in reaches dominated by knotweed than those dominated by trees. For a 9.5-month
- 18 monitoring period, there was 29 cm more erosion on banks that were also incised, and 9 cm more
- 19 erosion in banks with little incision. Turbidity responses to storm events were also higher (77 v 54
- 20 NTU) in reaches with knotweed, although this increase was found when the reach dominated by
- 21 knotweed was also incised. Thus, this study linked knotweed to increased erosion using multiple
- 22 methods.
- 23 Keywords: urban hydrology, bank pins, turbidity, knotweed
- 24

25 1. Introduction

26 1.2 Urban streams and sediment stressors

27 Degradation of urban streams and their receiving waters is common, resulting in a suite of 28 impairments dubbed "Urban Stream Syndrome" [1]. Waterways suffering from Urban Stream 29 Syndrome tend to have increased nutrient and contaminant concentrations, modified channel 30 morphology, reduced ecological function, and a hydrograph with steeper ascending and descending 31 limbs. Sediment fluxes change in urban systems because of the increased flow rates, and erosion and 32 deposition are one of the main stressors in urban streams. There is no defined amount of suspended 33 sediment associated with degraded streams as the turbidity in natural hydraulic systems can vary 34 significantly based on the density of the drainage basin, local geology, and the size of the stream or 35 river. 36 Turbidity is a measurement of the influence of suspended solids on an aqueous solution's

- 37 ability to transmit light. There are multiple concerns associated with turbidity, including, but not
- 38 limited to, decrease in light penetration, degradation of aquatic resources such as fish habitat,

39 sedimentation of receiving waters such as lakes and estuaries, and transport of other contaminants

40 [2, 3, 4, 5, 6]. Erosion can have major impacts on stream health by decreasing channel complexity,

41 disconnecting urban streams from their riparian zones, and increasing stream turbidity [5, 7,

42 8]. Sediment has been identified as one of the most significant pollutants entering the Chesapeake

43 Bay and legislation has been introduced to implement total maximum daily loads (TMDLs) to

44 control the amount of suspended sediment entering the estuary [9]. The Piedmont has been

45 identified as the single greatest sediment source into the Chesapeake Bay despite low relief and low,

46 long-term erosion rates [8, 9].

47 1.3 Measuring Fluvial Sediment Erosion

48 Precise measurements for bank erosion, sediment transport, and channel and floodplain 49 deposition to develop stream-scale sediment budgets are difficult to obtain. Sediment flux can also 50 be monitored by continuous data loggers monitoring turbidity. Sediment levels are often 51 associated with high flow events in response to precipitation. As such, regular sampling intervals 52 used in other water quality studies can miss elevated sediment concentrations during storm events 53 resulting in an underestimate of sediment load [10, 11, 12]. Skarbovik and Roseth [13] showed that 54 turbidity loggers were particularly successful at detecting peak concentration of sediment during 55 storm events. A site-specific calibration curve can be created to relate suspended sediment 56 concentration to the turbidity data. This allows turbidity to be used as a proxy for total suspended 57 sediment within the water column [10,14]. Turbidity loggers are not typically used on a reach scale, 58 instead focusing on catchment level changes in sediment load. Watershed scale monitoring is 59 especially useful in areas where increases in sediment are associated with non-point sources, an 60 issue that has been recognized for a number of pollutants since the early 1980s [5, 12, 15]. 61 Bank pins have been used to compare erosion in varying land-use which represented urban,

62 urbanizing, and agricultural watersheds as part of water quality assessment of sediment 63

contributions to Chesapeake Bay. These studies used bank pins to measure cut-bank erosion and 64

clay pads, or artificial horizons, installed to monitor floodplain deposition [9]. Measurements in 65

these catchments showed that the net site sediment budget was best predicted by the ratio of the

66 channel to floodplain width.

67 1.4 Riparian Zones and Sediment Flux

68 Riparian zones are broadly defined as semi-terrestrial areas that represent the interface of the 69 terrestrial and aquatic environment [16, 17, 18]. Riparian zones are generally dominated by woody 70 plants and are classified as shrub land or forest vegetation [16]. The relationship between vegetation 71 in the riparian zone and sediment dynamics is an important factor in stream geomorphology. When 72 trees are removed it results in extreme erosional potential as seen after changes in land management 73 in the America's after European settlement [19].

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74 The extent to which flora affects flow dynamics and stabilizes the banks is dependent on the 75 type of vegetation present as well as the density and depth of root systems. Vegetation can 76 decrease soil erosion by providing structural support through root development and increasing soil 77 cohesion by supplying organic material and influencing the soil moisture content [20]. One study, 78 [21], experimented on the effect of vegetation on erosion and found a 20,000-fold increase in 79 erosional resistance with an 18% increase in the volume of root mat in silty bank materials. 80 However, another study [22] noted that vegetation only protects against erosional processes if the 81 roots extend to the toe of the bank and a decrease in bank stability due to vegetation can cause mass 82 failure [16, 23]. In other words, vegetation on incised banks is not effective. If the ratio of rooting to 83 channel depth is small, then the erodibility of bank sediments and hydraulics of flow become more 84 important. Modeling efforts continued to increase the scientific community's understanding on the 85 influence of plant root systems on bank stability and erosion resistance [20, 24, 25]. 86 Riparian systems are particularly prone to colonization by invasive species. Stream corridors 87 facilitate invasion and propagation through transport and flooding while providing nutrients and 88 sediments with high levels of moisture [16, 18, 26]. Japanese knotweed (Reynoutria japonica) is 89 native to Japan, Taiwan and northern China [27, 28] and was introduced to Europe and America in 90 the mid-1800s as an ornamental plant. Considered one of the most successful plant invaders, it has 91 spread prolifically across both continents. The plant consists of bamboo-like stalks typically greater 92 than three meters in height and spreads predominately via mono-specific stands and aquatic 93 transport of rhizomes [27, 28, 29, 30, 31]. Rhizomes are defined as a laterally growing underground 94 stem that puts out shoots and roots.

95 Once established, knotweed is believed to promote bank erosion due to the shallow nature of 96 root systems when compared to riparian trees or shrubs. However, despite numerous references to 97 this phenomenon, there are few quantitative studies. One study, [27] shows an increase in sediment 98 load downstream of knotweed reaches after storm events but could not conclusively evaluate 99 whether this was associated with higher rates of erosion due to limited precipitation events.

100 *1.4 Objectives*

101 The objective of this study was to examine the influence of riparian zone characteristics on 102 erosion using both bank pins and turbidity loggers. The study site was the Tookany Creek, an 103 urban stream in the greater Philadelphia region. Vegetation type and the degree of incision were 104 used to classify riparian systems. We expected erosion rates would be higher in reaches with 105 Japanese knotweed, despite dense vegetation on banks, and we expected erosion rate to be higher in 106 steeply incised reaches (also known as disconnected from the riparian zone) in contrast to reaches 107 that have little or no incision which more readily allows bank overflow (connected reaches). 108 It is currently not well understood how localized riparian conditions, such as incision and 109 invasive species, influence the success of management practices implemented within the stream

110 channel. In addition, monitoring techniques used to examine sediment dynamics do not typically

111 measure the effects of local conditions and best management practices over short spatial scales. This

- 112 research is part of an effort initiated by the William Penn Foundation to improve water quality in
- 113 suburban Philadelphia watersheds through implementation of stormwater management practices
- 114 including riparian reconstruction. By understanding how riparian conditions can influence
- 115 sediment dynamics, more targeted placement of best management practices and stream
- 116 reconstruction can be implemented. This study also addresses the lack of quantitative analyses
- 117 concerning increased erosion along knotweed dominated stream corridors.

118 2. Materials and Methods

119 2.1 Study Area

120 The Tookany/Tacony-Frankford (TTF) Creek flows from Abington Township, Montgomery 121 County and through Philadelphia County before discharging into the Delaware River. Based on 122 calculations done in ArcGIS, the total TTF watershed is 93.6 km² with 41.6 km² (44%) of the 123 watershed designated the Tookany Watershed, upstream of the Philadelphia County line. In 124 addition to the main stem, there are five main tributaries in the headwater region. Together, their 125 total length is 34.2 km [32]. The average discharge of the Tookany upon entering Philadelphia 126 County is 0.74m³/s, based on data obtained from the USGS stream gauge 01467086. The Tookany 127 watershed mostly overlies mica schist. 128 Highly urbanized, the watershed is predominately low to medium level development with 129 51% of the land use classified as low intensity residential composed of single family detached 130 residential housing. Deciduous forest is the second most common land cover at 16% [32]. The 2008 131 stormwater management plan cites flooding, erosion, sedimentation, groundwater impacts, and 132 pollution as major issues associated with stormwater in the watershed [32].

133

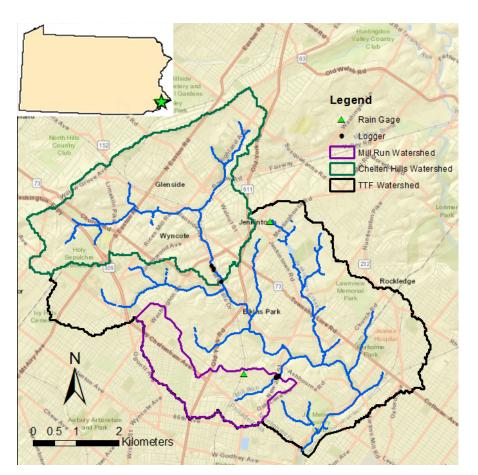


Figure 1. Tookany Watershed and PA state map inset: The Tookany Watershed (black) is 41.6 km² and is primarily classified as low intensity residential development. Two sub watersheds, Chelten Hills (green) and Mill Run (purple) were monitored using YSI loggers. There were 8 logger locations in the watershed.

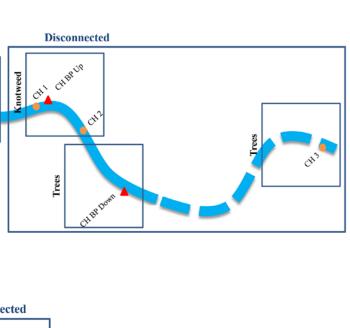
135Two reaches were chosen for monitoring within the Tookany Watershed (Figure 1). One136reach paralleled Chelten Hills Rd (CH), located on the main stem of the Tookany Creek. The137drainage area upstream of the monitoring locations represents 33% of the total Tookany Watershed.138Mill Run (MR) is a tributary to the Tookany Creek and monitoring locations represents 93% of the139tributary drainage area and 10% of the total Tookany Watershed. Each site had four turbidity140loggers (CH0 through CH3 and MR0 through MR3) and two bank pin monitoring locations (Figure1412) as described in the next section.

142 Discharge was measured along each reach near loggers 0 and 2. The discharge changed 143 little between logger locations, with the difference being within measurement error. The discharge 144 at Chelten Hills was 0.11 m³/s at baseflow and at Mill Run it was 0.02 m³/s at baseflow. Connected

Approximately 40 m

А

CHO



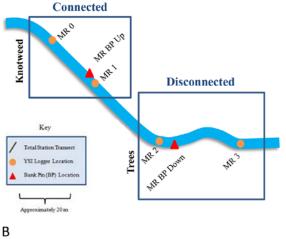


Figure 2. Sketch map of study reaches at Chelten Hills (A) and Mill Run (B) showing the locations and descriptions of each reach. Solid lines are drawn approximately to scale, dashed line indicate change in mapped distance. Stream flow is from station 0 to 3. Bank incision and riparian cover defined by boxed areas.

145 2.2 Monitoring Erosion

Bank pins and turbidity loggers were used to monitor erosion directly and indirectly,
respectively. These methods provide data on temporal and spatial scales within each monitored
reach. Four turbidity loggers and two bank pin locations were instrumented in each of the two
study reaches.

Bank pins are a method of monitoring erosion or deposition through the installation andmeasurement of exposed rods. By measuring the amount of pin exposed it is possible to determine

- 152 the extent of change as well as the rate of change. Bank pins were installed at two locations with
- 153 varying riparian characteristics, specifically vegetation and the degree of incision, at each study
- 154 reach (Figure 2). At Chelten Hills, the bank pins represent erosion rates in a disconnected riparian
- 155 section dominated by knotweed as well as a disconnected section with trees present. At Mill Run
- 156 bank pins where installed in a connected reach dominated by knotweed and a disconnected reach
- 157 dominated by trees. Along the connected reach, a small, less than 0.5 m bank allowed for
- 158 installation, but the reach was classified as connected due to the interaction between the stream
- 159 surface and both the floodplain and root structures.
- 160 Four erosion pins (0.5 m long garden stakes) were installed perpendicular to the cut bank
- 161 face. The pins were installed in a diamond pattern (Figure 3). There was approximately 0.5 m
- 162 between each pin with the bottom pin located just above the waterline present on the date of
- 163 installation.
- 164

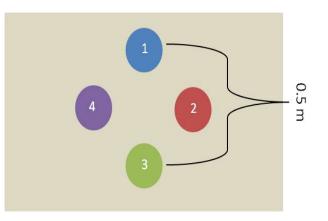


Figure 3. Bank pin installation diagram. Four bank pins were installed horizontally into the stream bank in a diamond pattern at each site. The length of the exposed pin was used to calculate erosion.

165 The initial length of the exposed pin was recorded and used to establish a starting position of

- 166 the bank. Subsequent measurements of the exposed pin were used to calculate how much erosion
- 167 had occurred at each pin location. Bank pins that experienced substantial erosion were reset to 15
- 168 cm when in danger of being eroded out of the bank. This occurred twice at the upstream Chelten
- 169 Hills monitoring point and once at the downstream Mill Run location. The measurement interval
- 170 ranged from five days to a month. At each site the change in bank position was analyzed
- 171 individually, then averaged. The average of the four bank pins was used to compare different
- 172 reaches as the rates were similar within sites [33].
- 173 Turbidity loggers (YSI OMS 600) were installed at four sites on the main stem near Chelten Hills
- 174 Road and four sites on Mill Run from May 2015 through March 2016. The loggers were numbered 0
- 175 through 3 from upstream to downstream. All loggers were calibrated prior to deployment. The
- 176 YSI loggers use an optical sensor to record turbidity in nephelometric turbidity unit (NTU). In

177 addition to turbidity loggers, water level loggers (Onset HOBO U20-04) were installed at each

178 location to relate water level to turbidity.

179 Loggers were installed near the banks. Equipment installed in meanders were placed on the 180 cut bank side. Loggers CH 0 through CH 3 at Chelten Hills were installed in the middle of a 181 connected native trees riparian zone, as well as the beginning and end of a highly disconnected 182 riparian zone with knotweed. There were 80 meters between loggers CH 0 and CH 1 and 30 meters 183 between loggers CH 1 and CH 2. The distance between the CH 2 and CH 3 was approximately 300 184 meters. CH 3 was placed for a longitudinal study to examine spatial variability. At Mill Run MR 0 185 and MR 1 were installed at the beginning and end of the connected reach dominated by knotweed 186 and MR 2 and MR 3 were installed at the beginning and end of a disconnected reach with trees. 187 This downstream reach was separated from the upstream reach by both large woody debris and a 188 stormwater inlet channel. There was approximately 30 meters between each logger.

189 Turbidity measurements were not directly related to total suspended sediment as no sediment 190 calibration curve was created for the logger locations. However, a relationship between total 191 suspended sediment and turbidity is typically linear [6, 10, 11, 13, 34, 35] and was likely similar 192 among sites given the sediment characteristics.

193 2.3 Streambed and bank characterization

194 Monitoring locations were chosen based on accessibility, vegetation type, and degree of 195 incision, and the habitats were classified based on these characteristics. Riparian type was 196 determined based on the presence of knotweed versus trees versus a mixture. Banks were 197 classified as incised or disconnected from its floodplain if roots did not interact with the water. 198 Connected reaches showed interaction between the stream surface and both the floodplain and root 199 structures. These reaches could have up to 0.5 m of incision. Highly disconnected reaches showed 200 at least 2 m of incision, which indicates longer term erosion, while reaches termed disconnected had 201 between 0.5 and 1.5 m of incision. The geomorphic position and sediment were characterized at each 202 logger location. The geomorphic position was characterized in terms of pool, riffle, or run and 203 degree of incision.

204 The streambed was characterized for embeddedness to evaluate the influence of local 205 conditions on turbidity. Embeddedness is defined as the extent that coarse substrates, such as 206 gravel and boulders are surrounded and covered by sand other fine sediments. The embeddedness 207 of the streambed was surveyed by evaluating how easily coarse grains on the surface of the 208 streambed could be dislodged. A 60 x 60 cm grid was constructed with markers placed every 5 cm to 209 ensure that measurement positions were unbiased. The grid was set up near the center of the stream, 210 directly adjacent to each turbidity logger except for CH3 which was omitted from the sediment 211 characterization study due to its distance from the other sites. A scale of 0 to 5 was used to 212 characterize the ease at which the grains were removed [36]. The scale represented increasing 213 embeddedness with a result of 0 representing a node where a coarse grain was present and could be

- 214 dislodged with no resistance. A result of 5 indicated that the grain could not be pulled out of the
- 215 streambed nor could it be moved at all (Table 1). If there were no coarse grains present at the grid
- 216 node, a notation of "S" was recorded to signify that only sand/loose sediment was present.
- 217
- 218 Table 1: Embeddedness scale
- 219

Scale	Description
S	Sand. loose sediment
0	Completely loose
1	Dislodges easily
2	Dislodges with little resistance
3	Dislodges with some difficulty
4	Can move around but cannot dislodge
5	Not removable, no movement

221 Streambed sediment texture was also measured. Samples were collected with a shovel from the 222 top 3-5 cm for sediment characterization on a gravel, sand, silt+clay trilinear plot. Fines are likely to

223 be removed by streamflow when using a shovel to collect samples from the bed, but the proportions

of the other sediment sizes are still commonly used to compare bedding characteristics. Sieves

ranging from 0.063 mm to 25.4 mm were used to separate samples into gravel, sand, and silt+clay.

Clasts larger than 25.4 mm were not included in the plots as their weight would significantly bias theresults.

228 2.4 Data Analysis

229 The turbidity response to storms was characterized by the peak value. Turbidity responses 230 generally had a rapid ascending limb followed by a gradual return to equilibrium conditions (Figure 231 4). The peaks were distinct for each storm. These turbidity peaks were compared to site 232 characteristics across different storm events to determine whether turbidity differed based on 233 vegetation, degree of incision, bed sediment grain size, or storm characteristics. Not all loggers at 234 each site recorded during storms due to battery failure or temporary sediment clogging, but the 235 number of events is noted in the analysis and the same storms were compared at each site. 236 Storms were characterized by total precipitation and intensity. To be classified as a storm, 237 precipitation needed to last over one hour. If there was a greater than 10-hour break in 238 precipitation a new storm event was declared as water level responses tended to recover within that 239 time period as based on visual inspection. Two rain gauges were used to identify precipitation 240 events. Data were downloaded from the Philadelphia Water Department rain gauge PWD_30 closer 241 to Mill Run and from a Villanova University rain gauge near Chelten Hills starting on July 28, 2015

242 (Figure 1). In addition, the water level logger data provided a measure of storm response. The 243 water level peaks were compared to turbidity peaks using linear regression 244 To determine the relationship among turbidity and riparian and streambed characteristics the 245 statistical software package, number crunching statistical software (NCSS), was used. When 246 analyzing two groups the Mann-Whitney U test was used. For example, the comparison of tree and 247 knotweed vegetations types were compared with this method. If greater than two groups were 248 being compared Kruskal-Wallis one-way analysis of variance (ANOVA) was used. Both analyses 249 were chosen as they could rank data that was not normally distributed and can be used on unequal 250 sample sizes. When only two categories were present for comparison T-tests were used, otherwise 251 ANOVA was used. Because several of the parameters used rank statistics rather than continuous 252 variation, multi-variate analysis was not applicable. Filters were applied to separate how 253 vegetation type, degree of incision, antecedent conditions, and seasonality influenced turbidity 254 levels. Antecedent conditions considered rain events within the last five days as wet conditions. If 255 no storms had occurred within five days conditions were considered dry. Seasonality was 256 examined using foliage as an indicator by splitting the dataset in mid-November. The amount and 257 intensity of precipitation were also compared between foliated and non-foliated conditions to ensure 258 that there was no change in storm size that might influence turbidity results. A p-value of 0.05 was

- 259 used as the threshold of determining statistical significance at the 95% confidence interval. Ranks
- 260 were used for parameters that did not vary continuously such as embeddedness and vegetation, and
- 261 for parameters with distinct groups such as incision, and grain size distribution.

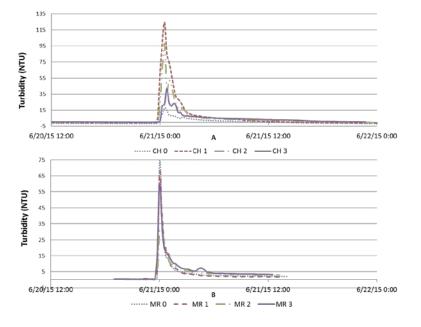


Figure 4. Example turbidity response to storm event. Turbidity response to storm on June 21, 2015 at Chelten Hills (A) and Mill Run (B).

262 **3. Results**

263 3.1 Logger site comparison

The vegetation along the banks included mature hardwood trees with visible root structure on the banks. Sometimes knotweed would be present in reaches with tress, but when the presence of tree roots was observed the section was labeled as tree-covered (Table 2). In other reaches, only knotweed was present. The disconnected banks were 1 to 1.5 m high, with highly incised reaches

showing at least 2 m of bank height. The stream width at Chelten Hills was 2 to 3 m, except CH0

- where it extended 4.75 m, resulting in a depth to width ratio (D:R) of 0.7 and 0.8 for highly
- disconnected banks and 0.3 for the connected bank (Table 2). Mill Run was generally wider, 6 to 7.5
- 271 m wide except at MR3 which narrowed to 4.5 m. The depth to width ratio was 0.2 except at MR3
- 272 where is was 0.4.
- The streambed sediments were characterized at each logger site except for CH3 (Table 2). Sand content in pools ranged from 29.4% to 92.6%, with the highest sand content appearing at Mill Run logger 2 (MR 2). At Mill Run the sediment collected from pools was found to have higher
- 276 proportions of sand to gravel than those sampled from the run. This pattern does not hold to
- proportions of sand to gravel than those sampled from the run. This pattern does not hold true atChelten Hills where the pool sample (CH 2) contains 16.0% less sand than the sample taken from the

Chelten Hills where the pool sample (CH 2) contains 16.0% less sand than the sample taken from therun (CH 0). The amount of gravel in runs ranged from 54.1 (CH 0) to 61.9% (MR 1). The greatest

- amount of gravel was found within riffles at both sites with CH 2 and MR 0 containing 74.3% and
- 280 71.7% gravel respectively.
- In both streams the embeddedness survey showed the majority of sediment (76 to 100%) was
 either sand/loose sediment (S) or loose grains (0-2) and only two sites showed more than 20%
- 283 embedded grains of some degree (Table 2). The two sites with strongly embedded grains were in
- riffles, and one of the runs has somewhat embedded grains.
- 285

Table 2: Site characteristics for turbidity logger locations. Loggers in bold showed statistically

- 287 different response from some of its neighbors.
- 288

Logger	Connectivity	D:R	Vegetation	Grain size	Geomorph	Embeddedness
CH0	Connected	0.3	Trees	Gravel	Run	Loose 79%
CH1	Highly Disconn	0.8	Knotweed	Gravel	Riffle	Loose 92%
CH2	Highly Disconn	0.7	Knotweed	Gravel	Pool	Loose 100%
CH3	Disconnected		Trees		Run	
MR0	Connected	0.2	Knotweed	Gravel	Riffle	Loose 76%
MR1	Connected	0.2	Knotweed	Gravel	Run	Loose 98%
MR2	Disconnected	0.2	Trees	Sand	Pool	Loose 100%
MR3	Disconnected	0.4	Trees	Sandy gravel	Pool	Loose 99%

289 3.2 Bank Pins: Erosional Events

290 Bank pin measurements showed erosion occurred episodically, but with an overall trend that 291 differed for each site. For example, at CH BP Upstream located on a highly incised cut bank with 292 mixed tree and knotweed vegetation, the degree of erosion varied during erosional events with no 293 one event being recorded in all bank pins (Figure 5). The variation between the four pins illustrates 294 the limitations in using bank pins to estimate erosion rates over larger scales. An extended change in 295 erosion rate occurred over the winter in three of bank pins locations, including CH BP Upstream. At 296 the CH BP Downstream, little erosion was observed. Because of this change in rate, the average 297 erosion was calculated separately for summer and early fall versus later fall and winter. The degree 298 of erosion at each of these pins was higher than at the other sites; thus, the rates for each site were 299 compared as an average rate.

300 The trend in average erosion rates for each site was the same across both seasons (Figure 6).

301 The site with disconnected banks and knotweed (CH BP Upstream) had the highest rate followed by 302

knotweed on a connected bank (MR BP Upstream). The lowest rates were observed at the two sites

303 with trees, both with disconnected banks (MR and CH BP Downstream).

304

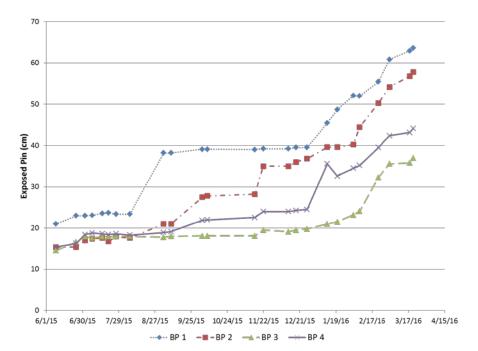


Figure 5. Chelten Hills upstream bank pin results. The change in the exposed bank pin as a measure of erosion. Increases in the rate of erosion are designated as erosional events.

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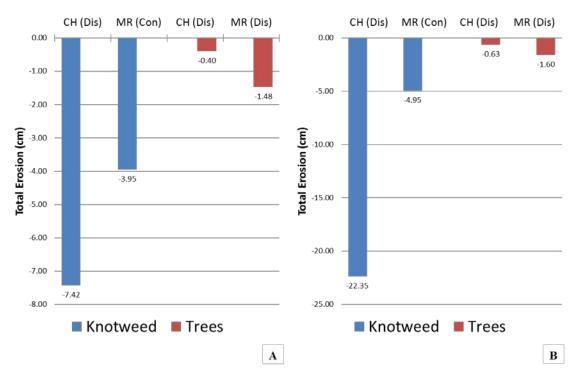


Figure 6. Total erosion at bank pin locations. The total erosion at each location from July to October 2015 (A) and October 2015 to March 2016 (B). The greatest amount of erosion occurred at the upstream highly disconnected reach containing mixed knotweed (CH BP Up) followed by a connected reach with knotweed (MR BP up). The least amount of erosion occurred at the downstream disconnected reach containing trees (CH BP Down) followed by a disconnected reach with trees (MR BP Down).

307

308 3.3 Turbidity Response to Storm Events

309

Continuous turbidity loggers recorded increases in turbidity for 43 storm events at Chelten
Hills and Mill Run. Logger response was compared for logger location, with and without foliage,
degree of incision, riparian cover, geomorphic position, and streambed grain size. Results are
presented as box plots and evaluated for statistically significant differences.
Linear regression between the water level peaks and the turbidity peaks showed the expected

Linear regression between the water level peaks and the turbidity peaks showed the expected relationship between higher storm flow and higher turbidity with one exception (Table 3). Most of the R² were 0.6 or better, but at MR2 the regression was only 0.25. The lower correlation at this site was attributed to woody debris upstream which caught debris; periodic release of debris and associated sediment could have lead to a noisier turbidity signal at this site. However, neither the mean nor the standard deviation was higher at MR2, so the woody debris did not contribute significantly to the sediment loading downstream. Although higher velocity associated with larger storms is expected to increase turbidity, the correlations at these sites suggest that other factors are

324 Table 3: Linear regression between water level peak and turbidity peak

0
R ²
0.60
0.84
0.77
0.59
0.73
0.76
0.25
0.75

325

326 Comparison of box plots showed that the turbidity response increased between loggers CH0

327 and CH1, which represented the transition from the connected with trees to a disconnected reach

328 with knotweed (Figure 7). The change in turbidity between these loggers was found to be

329 statistically significant (P-Value 0.006) using Kruskal-Wallis One-Way ANOVA with a difference in

330 mean turbidity of 81.7 NTUs. No significant difference (P-value 0.665) was found between loggers

331 CH1 and CH2 at the 95% confidence interval. CH3 had a mean turbidity of 46.4 NTU lower than

that of CH2. However, because the turbidity at CH3 had a high variability, this difference in mean

turbidity was not statistically significant (P-value 0.127) at the 95% confidence interval.

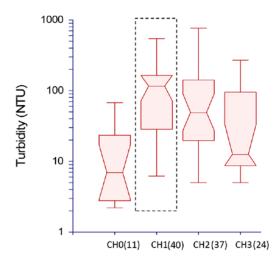


Figure 7. Turbidity response at Chelten Hills by logger. The turbidity response for each logger was plotted and analyzed with ANOVA. Logger CH0 was found to have a statistically lower turbidity response than loggers CH1 and CH2 but no statistical difference from other loggers. Numbers in parentheses indicate the number of storms included in the analysis.

335 The range of turbidity response at Mill Run was somewhat less than that of Chelten Hills. The 336 mean turbidity reading of logger MR2 was lower than the other loggers (42.3 NTUs lower than 337 logger MR0, 49.4 NTUs lower than logger MR1, and 25.0 NTUs lower than logger MR3). It was also 338 found to be statistically different (P-value 0.009) from loggers MR0 and MR1, but not from MR3 339 (Figure 8). This logger was located at the transition from knotweed to trees, but also located on a 340 disconnected reach. Thus, the logger with statistically lower turbidity at Mill Run was located in a 341 reach with trees and the logger with statistically higher turbidity at Chelten Hills was located in a 342 reach with knotweed.

Neither the geomorphic position nor the grain size clearly distinguished the two logger sites with statistically lower turbidity response, CH0 and MR2 (Table 2). CH0 is located in a run, but CH3 was also and it was not statistically different than the other Chelten Hills loggers. MR2 was one of two loggers located in pools. Thus, the two loggers with lower turbidity had different geomorphic position. CH0 had similar grain size distribution to most of the other loggers (gravely sand). MR2 was the only logger in sand, but the finer grain size would be expected to lead to higher, not lower turbidity.

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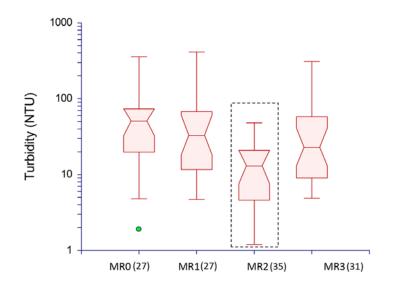


Figure 8. Turbidity response at Mill Run by logger. The turbidity response for each logger was plotted and analyzed with ANOVA. Logger MR2 was found to have a statistically different from loggers MR1 and MR2. Numbers in parentheses indicate the number of storms included in the analysis.

351 Riparian vegetation was compared at each site by examining the effect of knotweed on

352 turbidity in contrast to sites with mature trees. At Chelten Hills, sites monitored in tree-dominated

353 riparian zones experienced lower turbidity responses than knotweed dominated reaches (Figure 9).

354 Knotweed-dominated reaches had a mean value of 91.3 NTU and tree-dominated reaches had a

- 355 mean value of 53.3 NTU, a significant difference at the 95% confidence interval (P-value 0.02). Mill
- 356 Run had turbidity responses in knotweed dominated reaches with a mean value of 55.3 NTU and in

- 357 tree-dominated zones showed a mean value of 54.9 NTU. There was no statistical difference
- 358 (p-value 0.205) between the turbidity responses of different vegetation types at Mill Run. One
- 359 difference between the sites is that the knotweed section at CH is highly disconnected and it is
- 360 connected at MR (Figure 2). These results suggested that while the presence of knotweed increases
- 361 turbidity, the increase is not significant unless the stream is also incised.
- 362

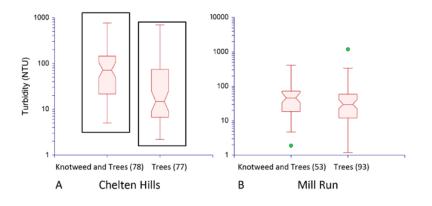


Figure 9. Turbidity response by vegetation. Outlined box plots indicate a statistical difference. When loggers were separated by vegetation type there was a statistical difference between mixed vegetation and trees only at Chelten Hills (A) but no statistical difference between the vegetation types at Mill Run (B). The logger site with trees at Chelten Hills was also highly incised.

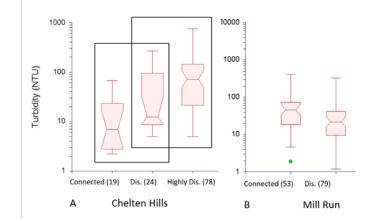


Figure 10. Turbidity response by degree of incision. Outlined box plots indicate a statistical difference. When loggers were separated by degree of incision there was a statistical difference between connected and highly disconnected reaches at Chelten Hills (A) and there was no statistical difference between the connected and disconnected reaches at Mill Run (B).

364 The degree of connectivity was also considered as a factor influencing turbidity response, with 365 the highly disconnected reaches showing greater turbidity. At Chelten Hills, a statistical difference 366 (P-value 0.016) was found between the connected and highly disconnected reaches but not between 367 the connected and disconnected reaches or between the disconnected and highly disconnected 368 reaches (Figure 10a). The mean values for connected, disconnected, and highly disconnected were 369 13.5 NTU, 40.5 NTU, and 86.4 NTU respectively. When loggers at Mill Run were separated by the 370 degree of incision, the mean response in the connected reach was found to be elevated (55.2 NTU) as 371 compared to the mean response of 32.0 NTU in the disconnected reach, although this difference was 372 not significant (Figure 10b). 373 The antecedent conditions were also examined to determine if dry conditions yielded different 374 turbidity responses than wet conditions. Dry conditions were defined as a period with no 375 precipitation for five or more days. During wet conditions the difference in turbidity between 376 degrees of incision was statistically significant at both Chelten Hills (P-value 0.010) and Mill Run 377 (P-value 0.005). There was no statistical difference between varying incision types at Chelten Hills 378 and Mill Run in dry conditions (P-values of 0.521 and 0.063 respectively). 379 Two statistical comparisons were made to evaluate the influence of foliage. One test evaluated 380 whether the turbidity response changed at a site when it has foliage versus without foliage. No 381 statistical difference between full and no foliage was found at either at Chelten Hills (P-value 0.705) 382 or Mill Run (P-value 0.519) at the 95% confidence interval (Figure 11). However, at both sites the 383 mean and standard deviation was found to increase when no foliage was present. Furthermore, the 384 statistical difference between knotweed and trees changed with foliage versus without foliage. At 385 both Chelten Hills and Mill Run there was a significant difference (P-values 0.018 and 0.012 386 respectively) between knotweed and trees when foliage was present. In contrast, there was no 387 difference at either site when foliage was not present (P-value 0.335 and 0.106). These differences 388 further emphasize the influence of knotweed on erosion, particularly in the winter months. This 389 increase in turbidity was not associated with corresponding increases in precipitation as neither the 390 amount of precipitation or intensity of precipitation increased during the winter months. 391

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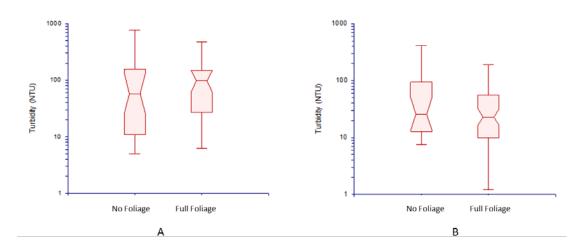


Figure 11. Turbidity Response to Change in Foliage: Turbidity data at loggers were analyzed to determine the influence of seasonality, based on the presence of foliage, at Chelten Hills (A) and Mill Run (B). While the mean turbidity does increase at both sites when no foliage was present, this change was not found to be significant at either site.

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401 4. Discussion

402 4.1 Summary of findings

Erosion was measured directly using bank pins and indirectly using turbidity loggers.
Seasonality appeared to influence erosion rates as an increase in erosion can be seen in bank pins
after December 2015. This compliments turbidity results which, while not statistically significant,
show an increase in the mean turbidity response when no foliage was present.

407 Riparian characteristics were found to influence erosion rates. For instance, turbidity

408 increased significantly (P-value 0.006) from CH 0 to CH 1 upon entering the disconnected reach

409 dominated by knotweed. Statistical analysis (Figure 7) showed vegetation to be the dominant

410 control on turbidity response. However, when comparing the turbidity response between

411 vegetation types at Mill Run, there was not a significant difference. Therefore, while vegetation

412 appears to be the dominant control, the degree of incision impacts whether the increase in erosion

413 associated with prevalent knotweed will result in a significant increase in turbidity. Bank pins also

414 indicated that there were higher rates of erosion in areas with knotweed, although the largest

415 differences were observed for reaches that were also incised.

416 4.2 Implications for bank management

417 Reconstruction of the entire riparian system is impractical due to the nature of property lines, 418 infrastructure, and cost. By determining which stream characteristics are beneficial and which 419 result in degradation of water quality, measurements on the effects of smaller riparian reaches, such 420 as those conducted for this study, can help promote effective practices [12, 15, 37]. Without 421 consideration of these local characteristics, management practices have a higher risk of failing 422 While many researchers assume the presence of knotweed will lead to higher rates of erosion, 423 previous studies quantifying that assumption are not reflected in the literature. One previous 424 study that provided quantitative analysis focused on rural catchments, which saw an increase in 425 sediment load downstream of knotweed reaches after storm events [27]. However, they not could 426 conclusively evaluate whether this was associated with higher rates of erosion due to limited 427 precipitation events. Other work has used models to predict effects of riparian vegetation [25]. Our 428 research provides quantifiable turbidity responses to storm events in reaches dominated by 429 knotweed within an urban setting. Our study suggests that while the presence of knotweed 430 increases erosion, the corresponding increase in turbidity is only statistically significant in 431 disconnected riparian zones. Therefore, if the goal of riparian restoration is to reduce sediment 432 loads, replacing Japanese knotweed with native vegetation by removing vegetation from the banks 433 is not necessarily effective where incision is minimal. Other management techniques that can be 434 used include, but are not limited to, those that promote in-stream deposition and bank stabilization 435 with rip rap or other engineering structures. Furthermore, if this management technique is used, it 436 is important that cleared riparian zones be maintained. The knotweed at Mill Run was removed as 437 part of a community riparian restoration but was recolonized within ten years. This suggests that the 438 removal of invasives might not be the best long-term solution if there are minimal funds for 439 continued monitoring and maintenance. Maintenance is also needed during the time it takes trees 440 planted for riparian restoration to mature, in order to prevent incision of the bank, loss of trees due 441 to disease, and invasion of exotic species such as Japanese knotweed.

442 4.3 Implications for sediment budgets

443 In streams where sediment erosion creates impairment, site and catchment scale sediment

444 budgets can be useful. Estimating erosion rates can be difficult, with multiple years of data

445 collection necessary [3, 11, 15, 38]. Once created, sediment budgets can be used to identify areas

446 that should receive direct management.

Sediment erosion calculations in the Tookany Watershed provided evidence for heterogeneity
based on specific riparian type in the catchment. Extrapolation to other reaches within the stream
could then be used to create a catchment scale budget, based on the influence of specific, small-scale

450	characteristics on erosion. However, this method requires an assumption that the data can be
451	projected to other locations within the catchment. This study had only two bank pin location per
452	reach. More comparisons of erosion rates between similar reaches would be useful to establish
453	variability before using this type of extrapolation. While the sample size within our study was not
454	large enough to provide a sediment budget, preliminary results indicate that bank pin monitoring
455	and turbidity data can be used to evaluate reach scale variations.
456	Small-scale studies like those in the Tookany Creek help to show the erosion variability within
457	streams. This is especially true in urban watersheds where increased discharge from impervious
458	surfaces, development adjacent to the stream corridor, and invasive species often culminate in
459	highly impaired waterways. Understanding sediment dynamics to decrease turbidity and
460	sediment levels entering water bodies is needed for improvement of stream health and water quality
461	and monitoring of streams is necessary to design and implement appropriate management practices.
462	
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466	
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470	Conflicts of Interest: The authors declare no conflict of interest.
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